1	Effects of biochar on carbon and nitrogen fluxes in boreal forest soil
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23 Abstract

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Background and aims The addition of biochar to soil may offer a chance to mitigate climate change
by increasing soil carbon stocks, improving soil fertility and enhancing plant growth. The impacts of
biochar in cold environments with limited microbial activity are still poorly known.

Methods In order to understand to what extent different types and application rates of biochar affect carbon (C) and nitrogen (N) fluxes in boreal forests, we conducted a field experiment where two different spruce biochars (pyrolysis temperatures 500°C and 650°C) were applied at the rate of 0, 5 and 10 t ha⁻¹ to *Pinus sylvestris* forests in Finland.

Results During the second summer after treatment, soil CO_2 effluxes showed no clear response to biochar addition. Only in June, the 10 t ha⁻¹ biochar (650°C) plots had significantly higher CO_2 effluxes compared to the control plots. The pyrolysis temperature of biochar did not affect soil CO_2 effluxes. Soil pH increased in the plots receiving 10 t ha⁻¹ biochar additions. Biochar treatments had no significant effect on soil microbial biomass and biological N fixation. Nitrogen mineralization rates in the organic layer tended to increase with the amount of biochar, but no statistically significant effect was detected.

39 *Conclusions* The results suggest that wood biochar amendment rates of 5-10 t ha⁻¹ to boreal forest 40 soil do not cause large or long-term changes in soil CO₂ effluxes or reduction in native soil C stocks. 41 Furthermore, the results imply that biochar does not adversely affect soil microbial biomass or key N 42 cycling processes in boreal xeric forests, at least within this time frame. Thus, it seems that biochar 43 is a promising tool to mitigate climate change and sequester additional C in boreal forest soils.

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45 Key words: Biochar; biological nitrogen fixation; microbial biomass; nitrogen mineralization;
46 nitrification; soil respiration

47 Introduction

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49 Biochar is formed by heating organic material under low oxygen concentrations in a process known 50 as pyrolysis (Lehmann and Joseph 2012). The addition of biochar to soil is a potential tool for carbon 51 (C) sequestration and climate change mitigation because biochar is enriched in C and recalcitrant to 52 decomposition in comparison to the original biomass (Woolf et al. 2010; Gurwick et al. 2013). Biochar can also act as a soil conditioner enhancing plant growth by increasing soil microbial activity, 53 54 water holding capacity, cation exchange capacity and pH (Lehmann and Joseph 2012; Robertson et 55 al. 2012; Biederman and Harpole 2013; Thomas and Gale 2015). However, these changes in soil 56 chemical and physical properties may increase microbial biomass, microbial activity and the 57 decomposition of soil organic matter (Lehmann and Joseph 2012). Moreover, the labile C fractions 58 of biochar may accelerate the decomposition of old soil organic matter through the priming effect 59 (Cross and Sohi 2011; Zimmerman et al. 2011; Fang et al. 2015; Wang et al. 2015). In addition, 60 biochar may affect the chemistry of phenolic compounds which commonly inhibit the decomposition 61 of soil organic matter in boreal forest soils. Fire-derived charcoal have been found to adsorb phenolic 62 compounds and to accelerate organic matter decomposition in boreal forests (Zackrisson et al. 1996; Wardle et al. 1998, 2008). Accelerated decomposition of native soil C increases soil CO₂ emissions 63 and reduces the soil C stocks, which is contradicting the idea of C sequestration. 64

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The impacts of biochar addition on soil processes have been variable and are dependent on the pyrolysis temperature and the feedstock of biochar (Spokas and Reicosky 2009; Ameloot et al. 2013; Biederman and Harpole 2013; Lei and Zhang 2013; Stewart et al. 2013) soil properties (Kolb et al. 2009; Spokas and Reicosky 2009), vegetation and local environmental and climatic conditions (He et al. 2017). Previous studies have mainly been conducted on agricultural soils in tropical and temperate regions, and very little information exists about the stability of biochar in the soil and the effects of biochar additions on C and nutrient cycling in forests, especially in the boreal zone (Liu et al. 2015; Bruckman et al. 2016). The use of forest biomass as an energy source has increased in Europe (Helmisaari et al. 2014). Instead of traditional burning, part of the forest biomass could be converted to biochar, which can be incorporated back into soil, where it helps to improve the sustainability of bioenergy harvesting if part of C and nutrients were recycled back to the forests and if biochar acts as a soil amendment.

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79 In boreal forests, most of the soil nitrogen (N) is in organic form, N mineralization rates are low and 80 tree growth is N-limited (Sponseller et al. 2016). The mineralization of N can be accelerated if biochar 81 stimulates soil organic matter decomposition which, in turn, may have a positive feedback on 82 ecosystem net primary production and CO₂ fixation. Biochar application has been shown to increase 83 net N mineralization and nitrification rates (Ameloot et al. 2015; Case et al. 2015; Gundale et al. 84 2015) which has been attributed to increased soil pH, enhanced microbial growth and activity and the 85 sorption of phenols and terpenes onto biochar (Clough and Condron 2010; Lehmann et al. 2011). 86 Polyphenolics and terpenes inhibit nitrification and net N mineralization by decreasing the activity of 87 enzymes involved in N cycling (Adamczyk et al. 2015, 2017). Wildfire-produced charcoal has been 88 found to adsorb phenols, and to increase net N mineralization and nitrification in forest soils 89 (Zackrisson et al. 1996; Wardle et al. 1998; DeLuca et al. 2006; Ball et al. 2010). Biochar may thus 90 serve as an important soil amendment, and it could be possibly used for mimicking the effects of firederived charcoal in Finland, where forest fires are effectively controlled (total area of forest fires is 91 only 300-1000 ha⁻¹ yr⁻¹) and forest soils contain high amounts of phenolic compounds. On the other 92 93 hand, the reduction of N mineralization and increased N immobilization may occur when biochar 94 compounds with a high C:N ratio are microbially degraded (Bruun et al. 2012; Dempster et al. 2012; 95 Prommer et al. 2014) and due to the adsorption of NH₄⁺ or NO₃⁻ onto the biochar surface (Clough 96 and Condron 2010).

98 Many boreal forests receive low amounts of N deposition and biological N fixation contributes 99 significantly to N input in these ecosystems (Granhall and Lindberg 1980; DeLuca et al. 2002; 100 Sponseller et al. 2016). Feather mosses that support epiphytic cyanobacteria represent the primary 101 source of biological N-fixation in boreal coniferous forests (Zackrisson et al. 2004), but there are also 102 free-living N-fixing bacteria in forest soils (Granhall and Lindberg 1980; Limmer and Drake 1996). 103 The influence of biochar amendment on N-fixation in boreal forests is not yet known. Biochar may 104 affect the magnitude of biological N-fixation by changing the biomass and species composition of 105 mosses (Zackrisson et al. 2004). Increased soil pH and more favourable soil moisture conditions after 106 biochar addition may enhance N-fixation (Nohrstedt 1985; Limmer and Drake 1996) whereas 107 increased availability of inorganic N may have a suppressing effect (Zackrisson et al. 2004; DeLuca 108 et al. 2007).

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110 The purpose of our study was to determine whether biochar additions increase soil pH, soil microbial 111 biomass and N transformations (net N mineralization, ammonification and nitrification) in boreal 112 forest soil. Additionally, we examined whether biochar affects soil CO₂ fluxes and biological Nfixation rates. We hypothesized that biochar amendment will increase soil pH and microbial biomass, 113 resulting in increased soil respiration, N-mineralization, nitrification and N-fixation. We also 114 115 hypothesize that these increases will occur to a greater extent at higher biochar amounts. The effects 116 of biochar on soil C and N fluxes were studied in the second year after the treatment. Generally 117 biochar causes at least a short-term limited positive priming effect (Bruckman et al. 2015; Mitchell 118 et al. 2015; Page-Dumroese et al. 2017), but the longer-term field experiments about the impacts of 119 biochar in forest ecosystems are rare. Biochar increased soil respiration in our study plots during the 120 first months after treatment (Palviainen et al. 2017a), and we wanted to know whether biochar 121 addition alters soil CO₂ effluxes for a longer term in boreal forest soil.

123 Materials and methods

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125 Study area

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The study area situates in southern Finland in Juupajoki (61° 48' N, 24° 18' E, 181 m a.s.l.) close to 127 Hyytiälä Forestry Field Station. The experiment was performed in young ~20-year-old Scots pine 128 129 (Pinus sylvestris L.) forest stands that were naturally regenerated from seed trees after clear-cutting. 130 The sites were nutrient poor xeric (Calluna) and sub-xeric (Vaccinium) forest site types (Cajander 1949). The mean height of trees was 5.0 m, diameter at breast height (1.3 m) was 4.9 cm, and the 131 number of trees (height > 1.3 m) was 4025 ha⁻¹. Understory vegetation is dominated by dwarf shrubs 132 133 (Vaccinium vitis-idaea L., Calluna vulgaris (L.) Hull., Empetrum nigrum L. and Vaccinium myrtillus 134 L.), mosses (Pleurozium schreberi (Brid.) Mitt. and Dicranum polysetum) and lichens (Cladina sp.). 135 The terrain is flat and the soil is a nutrient-poor, well-drained haplic podzol (IUSS Working Group 136 WRB, FAO 2015). The soil texture is coarse sand. The long-term (1981-2010) mean annual 137 temperature in the area is 3.5°C and annual precipitation is 700 mm (Pirinen et al. 2012). During the 138 experimental period in summer 2016, mean air temperature was 14.0°C in June and 16.0°C in July. 139 Precipitation was 124 mm in June and 119 mm in July in the year 2016.

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The experiment was set up as a replicated split plot experiment with four replicates (called whole plots) and five subplots (15 m × 15 m) within each whole plot. Whole plots were separated by a few hundred meters from each other and belonged to different forest stands to avoid pseudo-replication. The subplots were amended with biochar produced from Norway spruce (*Picea abies* (L.) H. Karst) wood chips at two different temperatures, at 500°C and at 650°C (manufactured by Sonnenerde GmbH, Riedlingsdorf, Austria). The biochar was produced by using the Pyreg process and the grain 147 size was 5-10 mm (Bruckman et al. 2015, Fig. 1). Both types of biochar were applied on the plots at two different amounts, 0.5 kg m⁻² and 1.0 kg m⁻². Thus, in each whole plot there were five treatments: 148 a control without biochar, 500°C biochar 0.5 kg m⁻², 500°C biochar 1.0 kg m⁻², 650°C biochar 0.5 kg 149 m⁻² and 650°C biochar 1.0 kg m⁻². There was a 10-meter buffer zone between each subplot. Biochar 150 151 was spread manually on the top of the organic layer during the last two weeks of May in 2015 (Fig. 152 1). Biochar was spread to the soil surface to avoid soil disturbance and damage to roots. The amounts of biochar correspond to 5 and 10 t ha⁻¹, which are typical and economically feasible biochar 153 154 application rates in forests (Bruckman et al. 2016). The added amounts of biochars were considerably higher than the amounts of charcoal, or black C (range 0-2220 kg ha⁻¹, mean 770 kg ha⁻¹) originated 155 by forest fires in Scandinavian boreal forests (Ohlson et al. 2009). 156

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158 Soil and biochar analyses

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160 Soil samples were collected from the organic layer and the upper 15 cm mineral soil layer using 161 stainless soil corer (diameter 5.5 cm) at nine locations in each subplot in mid-May in 2015 just before 162 biochar addition. The samples were dried (60°C, 24 h), sieved through a 2-mm sieve, and ground 163 before the analysis. Subsamples were taken for dry mass determination at 105°C. Soil particle size 164 distribution was determined by the laser diffraction (LS230, Coulter Corp., Miami, Florida, USA) 165 method (Table 1). The C and N concentrations of soil and biochars were analyzed with an elemental analyzer (Vario Max CN elemental analyser, Elementar Analysensysteme GmbH, Germany). The 166 167 loss on ignition (LOI) of biochars were determined by combusting samples at 550°C for 3 hours. The 168 concentrations of P, K, Ca, Mg, S, Fe, Al, Na, Cu, Mn, Ni, Si and Zn in biochar were determined 169 from HNO₃-H₂O₂ digestion by ICP atomic emission spectrophotometer (ARL 3580 OES, Fison 170 Instruments, Valencia, USA). Biochar pH was determined using a pH meter (PHM210, Radiometer 171 Analytical, France) on a 1:2.5 (v:v) biochar /water solution and electric conductivity was measured by an electric meter (JENWAY 4010 Conductivity, TER Calibration Ltd., Wigan, UK). The
properties of biochars are presented in Table 2.

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175 Soil temperature and soil respiration measurements

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Soil temperature was measured continuously on all sample plots at three hours intervals with iButton
temperature sensors (Maxim Integrated, San Jose, California, U.S.A.), that were installed under the
organic layer. We interpolated hourly values from which we calculated daily mean temperatures for
each plot.

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Six polyvinyl chloride (PVC) collars (diameter 0.22 m) were installed permanently into the soil in each of the 20 subplots in the summer of 2015 for soil respiration measurements. Thus, there were 24 collars in each treatment and 120 collars in total. The lower edge of the collar was placed at 0.02 m depth in the mor layer above the rooting zone to avoid damaging the roots. The collars were sealed with a thin layer of sand placed around the collar. Ground vegetation inside the collars remained intact.

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Soil respiration i.e. CO₂ effluxes were measured with a closed chamber system consisting of an opaque cylindrical polycarbonate chamber (diameter 20 cm, height 30 cm), a CO₂ analyzer, sensor for relative humidity and temperature and a data logger (Kulmala et al. 2008; Pumpanen et al. 2015). The CO₂ concentration inside the chamber was recorded with a GMP343 diffusion type CO₂ probe (Vaisala Oy, Vantaa, Finland) at 5-second intervals and corrected automatically for humidity, temperature and pressure with a data recorder (MI70, Vaisala Oyj) using the readings from the temperature and humidity probe (HMP75, Vaisala Oyj) inside the chamber. Air pressure was measured daily at the nearby SMEAR II station (4 km away). During the measurements, air insidethe chamber was mixed continuously by a small fan.

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199 The chamber was placed onto the collars only during the measurements which lasted 4 minutes. Soil 200 respiration measurements were conducted with two chambers in two consecutive days in June and 201 July 2016 (i.e. 13 and 14 months after biochar addition). All collars were measured before noon to minimize daily temperature fluctuations. Air temperature during the soil respiration measurements 202 203 varied $\pm 0.7^{\circ}$ C in June and $\pm 1.0^{\circ}$ C in July, and the variation in soil temperatures was even smaller 204 (±0.3°C) indicating that temperature fluctuations during the measurements did not markedly affect the results. Headspace volume was corrected for the varying height of the collars. Soil temperature at 205 5 cm depth was measured by a dual input digital thermometer (Fluke-52-2, Fluke Corp.) 206 simultaneously near the collar. The CO₂ efflux was calculated as the slope of a linear regression of 207 208 CO₂ concentration in the chamber against time. Only measurements taken between 45 seconds and 3 209 minutes after the closure were included in the fitting.

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211 Nitrogen mineralization experiment

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Nine soil core samples (diameter 5.5 cm) were collected in November 2016 from the organic layer 213 214 and the upper 10 cm mineral soil layer from the control subplots and from the subplots where 650°C produced biochar were added 5 t ha⁻¹ and 10 t ha⁻¹, respectively. Soil samples were stored at $+5^{\circ}$ C in 215 216 plastic bags for a few days before further treatment. The nine soil samples from each subplot were 217 combined to give three composite samples per subplot (n= 12/treatment). To homogenize the soil 218 material, the samples were sieved through a 2-mm sieve. Nitrogen transformations were studied by 219 incubating 10 g of humus and 20 g of mineral soil in cork sealed 125-ml glass bottles in a climate 220 chamber (WEISS WK11 340, Weiss Klimatechnik GmbH, Germany) at constant temperature (15°C) 221 and moisture (soil moisture content adjusted to 60% of the water-holding capacity) for 42 days. At the start and at the end of the incubation, an analysis of inorganic N was performed to estimate net N 222 mineralization, ammonification and nitrification for the samples. Each soil sample was extracted with 223 40 ml of 1 M KCl for 2 h (ISO 14256–2: 2005). The KCl extracts were filtered through a 0.45-µm 224 225 filter and ammonium (NH₄-N) and nitrate (NO₃-N) concentrations were analyzed with a flowinjection ion analyzer (Lachat Quickchem 8000, Milwaukee, WI, USA). Initial concentrations of 226 (NH₄⁺-N) and (NO₃⁻-N) were subtracted from the corresponding post-incubation concentrations to 227 228 calculate the rates of net ammonification and nitrification. Net mineralized N was calculated from the 229 sum of (NH₄⁺-N) and (NO₃⁻-N) accumulated during the period of incubation. The incubated soil 230 samples were dried, ground with a mortar grinder (Retsch RMO Mortar Grinder, Retsch GmbH, 231 Germany) and their C and N concentrations were measured with an elemental analyser (Vario Max 232 CN, Elementar Analysensysteme GmbH, Germany). A subsample was taken for dry mass 233 determination (105°C, 24 h). The formed inorganic N was expressed on organic matter basis (µg N, NO₃ or NH₄ g C⁻¹ d⁻¹). 234

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Soil pH was measured from separate samples by mixing 10 ml of soil with 25 ml of deionized water.
The suspension pH (H₂O) was measured with a glass electrode (PHM210, Radiometer Analytical,
France) after 24 hours.

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240 Biological nitrogen fixation and moss biomass

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The samples containing mosses and organic layer were collected in May, June and July 2016 with a soil core cylinder (diameter 5.8 cm) from the control subplots and from the subplots where 650°C produced biochar were added 5 t ha⁻¹ and 10 t ha⁻¹, respectively. In total, 108 samples were collected for biological N fixation measurements (12 samples per treatment, 3 treatments and 3 sampling

times). Biological N fixation was estimated using acetylene reduction method (Hardy et al. 1968). 246 The samples included organic layer because in boreal forests N-fixation occurs both in the organic 247 layer and mosses (Granhall and Lindberg 1980; Limmer and Drake 1996). The whole samples were 248 249 placed in 500 ml glass jars with rubber septum caps, after which 10% of the volume of the jar was 250 evacuated using a gas-tight syringe (BD Plastipak 60, BOC Ohmeda, Helsingborg, Sweden) and 251 replaced with acetylene. The samples were incubated in an environmental chamber (WEISS WK11 340, Weiss Klimatechnik GmbH, Germany) with artificial light (LED Grow Light Spider 1) at 10°C 252 253 (samples collected in May), 15°C (samples collected in June) and 20°C (samples collected in July) for 24 hours. After incubation, a gas sample was taken from each jar by a 50-ml polypropylene syringe 254 (BD Plastipak 60, BOC Ohmeda, Helsingborg, Sweden), injected into a 12 ml exetainer vial (Labco 255 256 limited, Lampeter, UK) and the ethylene concentrations were analysed with a gas chromatograph 257 (HP6890) with flame ionization detector as described before (Leppänen et al. 2013). A commonly used ratio of 3 moles of reduced acetylene per mole of N fixed was used to calculate the mass of fixed 258 259 N (DeLuca et al. 2002). The biomass of different moss species was determined after drying the 260 samples at 60°C for 48 hours to see whether the biochar amendment affects the biomass of mosses, 261 and to explain possible differences in N-fixation rates between treatments.

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263 Soil r	nicrobial	biomass
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Twelve soil core (diameter 10.0 cm) samples per treatment were collected for microbial biomass C and N analysis both in June and July of 2016 from the organic layer from the control subplots and from the subplots where 650°C produced biochar was added. Root material was removed with tweezers, the samples were placed into 45 ml plastic tubes and stored in the freezer at -20°C. The samples were kept 7–10 days at + 5 °C before analysis. Samples were sieved through a 2-mm sieve, grinded (DeLonghi KG49) and a subsample was taken for dry mass determination (105 °C, 24 h). 271 Soil microbial biomass C and N were determined by a chloroform fumigation extraction method (Brookes et al. 1985; Vance et al. 1987). Three grams of soil from each sample was weighed, placed 272 273 into glass beakers and fumigated with 30 ml ethanol-free chloroform (CHCl₃) in a vacuum desiccator. 274 Another equivalent sample weighting three grams was placed in plastic bottles in another desiccator 275 as un-fumigated control samples. Both desiccators were kept at 25 °C in the dark for 24 hours. After 276 fumigation, 0.5 M potassium sulfate (K_2SO_4) (with the ratio of oven-dry basis soil: $K_2SO_4=1:20$) was used to extract the fumigated and un-fumigated samples. Then the samples were shaken at 200 rpm 277 278 for 1 hour and filtered using Whatman No.42 ashless filter papers. The filtrate was then used to 279 analyze the microbial C and N by a TOC-VCPH analyzer (Shimadzu Corp., Kyoto, Japan). Microbial 280 biomass C and N were calculated as the difference between fumigated and unfumigated samples and 281 the difference was divided by the soil-specific calibration factor which was 0.45 for C (Beck et al. 282 1997) and 0.54 for N (Brookes et al. 1985).

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284 Statistical analyses

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286 The effect of biochar amendment on soil pH, soil temperature, soil respiration, soil microbial biomass, 287 biological N fixation and N mineralization were analyzed with linear mixed model followed by 288 Fisher's least significant difference (LSD) test. Treatment was a fixed factor and plot was a random 289 factor. In the soil respiration analyses, the collar within the subplot was set as random factor. Data were checked for normality with the Shapiro-Wilk test and the recorded CO₂ effluxes were 290 291 logarithm-transformed. Differences were considered statistically significant when P was ≤ 0.05 . Statistical tests were performed using IBM SPSS version 23 (IBM Corp, Armonk, NY, USA). The 292 293 results of the statistical tests are presented in supplementary material.

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295 Results

297 Biochar characteristics

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Carbon concentrations were similar in both biochars, while the concentrations of N and other macronutrients tended to be lower in 650°C than in 500°C biochar (Table 2). Also C:N ratio was considerably higher in 650°C than 500°C biochar. Altogether, 3031 kg ha⁻¹ and 6061 kg ha⁻¹ of C were added to the soil along with 5 t and 10 t ha⁻¹ biochar treatments, respectively. These amounts correspond to 14% and 28% of soil C pools (organic layer and 0–15 cm mineral soil layer) in the study site (Table 1).

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306 Soil temperature and soil respiration

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308 Soil temperatures did not differ significantly among the treatments (Table 3). Treatment had significant effect on soil respiration in June (F= 3.978, P= 0.005) but not in July (F= 1.411, P=0.259). 309 310 Soil temperature as a covariate was not significant (June: F=0.852, P= 0.358, July: F=0.695, P=0.407) 311 and inclusion of this covariate in the analysis did not affect the results. In June, soil CO₂ efflux was significantly higher in plots where 650°C produced biochar was applied 10 t ha⁻¹ compared to control 312 and 5 t ha⁻¹ biochar treatments (Fig. 2). Both in June and July, 500°C biochar plots had higher soil 313 CO₂ effluxes in 10 t ha⁻¹ treatments compared to 5 t ha⁻¹ treatments (Fig. 2). The production 314 315 temperature of biochar did not have an effect on soil CO₂ fluxes, as there was no statistically significant difference in CO₂ effluxes between 500°C and 650°C biochar subplots in 5 t and 10 t ha⁻¹ 316 317 treatments.

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319 Soil microbial biomass, moss biomass and biological N fixation

The biochar treatments did not significantly influence soil microbial biomass C or soil microbial biomass N (Fig. 3). Microbial biomass C:N-ratio was significantly higher in 5 t ha⁻¹ biochar plots than in 10 t ha⁻¹ biochar plots in June, but microbial biomass C:N-ratios did not differed between treatments in July (Fig. 3).

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The total biomass of mosses was similar between control and biochar treatments, but there were slight differences in species abundances because the biomass of *Pleurozium schreberi* was significantly higher in 5 t ha⁻¹ biochar plots than in 10 t ha⁻¹ biochar and control plots (Table 4).

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Biochar amendment had no significant effect on biological N fixation rate (Fig. 4). Nitrogen fixation rates were significantly higher at an incubation temperature of 20°C (P <0.001) but did not differ between 10°C and 15°C. The mean N fixation rates were 199, 233 and 439 μ g N m⁻² d⁻¹, at 10°C, 15°C and 20°C, respectively. By taking into account the average length of growing season (180 days) and mean air temperature (~15°C) during growing season in the study area, the measured N fixation rates correspond to 0.56, 0.43 and 0.58 kg ha⁻¹ yr⁻¹ in control, 5 t ha⁻¹ and 10 t ha⁻¹ biochar treatments, respectively.

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- 338 Soil pH and N transformations
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Soil pH in the organic layer and the upper 10 cm mineral soil layer was significantly higher (P <0.04) in 10 t ha⁻¹ treated biochar plots than in the control plots (Fig. 5). In the control plots, soil pH was 3.7 in the organic layer and 4.1 in the upper 10 cm mineral soil, whereas in 10 t ha⁻¹ treated biochar plots the respective values were 4.1 and 4.3.

The biochar treatments did not induce statistically significant effects on net N mineralization, ammonification or nitrification rates (Fig. 6). The average net N mineralization rates in the organic layer increased with the amount of biochar, being 0.95, 2.30 and 2.78 μ g N g C⁻¹ day⁻¹ in the control, 5 t ha⁻¹ biochar plots and 10 t ha⁻¹ biochar plots, respectively. However, this difference was not statistically significant (P >0.05) due to the high variation within each treatment. In the mineral soil, net N mineralization was small or N was immobilized. Net nitrification was also negligible.

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352 Discussion

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Few studies have investigated *in-situ* the effects of biochar addition on soil respiration in forest 354 355 ecosystems. There was no clear and consistent tendency towards increased soil CO₂ effluxes during the second summer after biochar addition. Only in June, the CO₂ effluxes were significantly (17%) 356 higher in 10 t ha⁻¹ 650°C produced biochar plots than in the control plots. Otherwise, there were no 357 358 differences in soil CO₂ effluxes between control and biochar treatments. Slightly increased soil CO₂ 359 effluxes after biochar addition may be observed due to the mineralization of labile C fractions of 360 biochar and/or biochar induced priming effects in the soil shortly after biochar amendment (Smith et al. 2010; Zimmerman 2011; Cross and Sohi 2011; Jones et al. 2011). Biochar may also indirectly 361 stimulate microbial activity by providing nutrients, offering a habitat because of its porous structure, 362 363 increasing soil pH and reducing the bioavailability of toxic compounds in soil through sorption (Steinbeiss et al. 2009; Lehmann et al. 2011; Lehmann and Joseph 2012; Hammer et al. 2014). In 364 addition, biochar may increase plant growth and root biomass (Lehmann et al. 2011; Robertson et al. 365 366 2012; Thomas and Gale 2015), which promotes root respiration and provides additional organic matter for decomposition. 367

Both increased and decreased C mineralization has been observed following biochar addition to 369 various types of soils (Cross and Sohi 2011; Zimmerman et al. 2011; Liu et al. 2015; Wang et al. 370 371 2015). Studies from temperate forests have reported short-term positive priming effects or unchanged soil respiration after biochar addition (Sackett et al. 2014; Bruckman et al. 2015). Gundale et al. 372 (2015) mixed 10 t ha⁻¹ biochar to boreal forest soil and did not find significant effect on soil 373 374 respiration. In general, the positive priming effects are observed in soils which have low C contents 375 (Zimmerman et al. 2011). Weak priming effects and moderate changes in CO₂ effluxes in boreal 376 forest soils after biochar addition may take place since boreal forest soils have high C content (DeLuca 377 and Boisvenue 2012).

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379 The responses of soil CO₂ effluxes depend also on feedstock characteristics, pyrolysis temperature and application rate (Zimmerman et al. 2011; He et al. 2017). In general, wood biochars increase soil 380 381 CO₂ effluxes to a lesser degree compared to other types of biochars, and soil CO₂ effluxes decline 382 with biochar pyrolysis temperature (Zimmerman et al. 2011; He et al. 2017). In the present study, 383 wood biochar, produced at relatively high temperatures, may be the reason for the small changes in 384 soil respiration. Furthermore, apparently moderate biochar amendments do not cause large increases 385 in soil respiration. For example, meta-analyses from croplands have showed that biochar increases soil CO₂ emissions significantly only at high (20–40 t ha⁻¹) amendment rates (Song et al. 2016; He et 386 387 al. 2017). However, our results also showed that there was tendency for higher soil CO₂ effluxes from 10 t ha⁻¹ plots than from 5 t ha⁻¹ plots, at least in 500°C biochar treatments (Fig. 2). Pyrolysis 388 389 temperature did not have an effect on soil CO₂ fluxes, although generally biochars produced at high 390 (> 600°C) temperatures are more recalcitrant than those produced at lower temperatures (Cross and 391 Sohi 2011; Ameloot et al. 2013; Fang et al. 2015) and they often cause negative priming effects in 392 the soil (Zimmerman et al. 2011; Song et al. 2016).

394 The differences in soil CO₂ efflux among our biochar and control plots were higher during the first 395 summer (Palviainen et al. 2017a) when compared to the second summer. Soil CO₂ effluxes at 10 t ha⁻ 396 ¹ biochar treatments (both 500°C and 650°C biochar treatments) were significantly higher compared 397 to the control throughout the first summer and this effect was attributed to warmer soils after biochar 398 application to the soil surface (Palviainen et al. 2017a). In the second summer, biochar largely 399 disappeared under the moss layer (Fig. 1d), and soil temperatures were similar among treatments (Table 3) which likely reduced the differences in soil CO₂ effluxes between biochar and control plots. 400 401 Studies from temperate forest soils have also indicated that increases in soil CO₂ effluxes after biochar 402 addition are transient and can be generally observed only during the first year (Bruckman et al. 2015; 403 Page-Dumroese et al. 2017). In the long-term, biochar addition may even decrease the rate of soil C 404 mineralization because the adsorption of organic matter and microbial extracellular enzymes to biochar slows down the decomposition (Cross and Sohi 2011; Jones et al. 2011; Zimmerman et al. 405 406 2011; Ameloot et al. 2013; Prayogo et al. 2014).

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408 In growing forests, biochar can only be applied on the soil surface where it may be prone to losses 409 caused by surface runoff and wind. Bruckman et al. (2016) have studied biochar particle movement on a forest floor that is very similar to our experiment, by using terrestrial laser scanning in 410 combination with a time-lapse photography. They used similar biochar as in this study (grain size, 411 412 feedstock material, pyrolysis process conditions and post-production procedures) and found that 413 particle movement is slight and occurs only during heavy precipitation events or strong winds shortly 414 after biochar application to soil. In this study, little if any biochar was lost from the area with wind 415 because the forest was quite dense and biochar particles submerged below the ground vegetation and between the mosses during spreading. The transportation of biochar away with the surface water flow 416 417 is also unlikely because the terrain is flat, soil is well-drained coarse sand and there were no heavy 418 rains during the experimental period. Furthermore, the biochar was not a powder but the particle size

419 was 5-10 mm (Fig. 1a). Bruckman et al. (2015) applied similar biochar as in this study to the soil 420 surface in a temperate forest, and they found that the litter layer contained more C as compared to the 421 control plots. This surplus of C equaled the amount which was applied, suggesting that surface applied 422 biochar effectively incorporates in the organic layer shortly after amendment at the given surface 423 properties and application rates (Bruckman et al. 2015).

424

There was no effect of biochar addition on soil microbial biomass. Previous studies have also found that biochar additions of 5 to 10 t ha⁻¹ did not have significant effect on microbial biomass in forest soils (Sackett et al. 2014; Gundale et al. 2015; Noyce et al. 2015). The null effect on microbial biomass may be due to the low biochar addition rate. The more pronounced shifts in the soil microbial biomass have been observed with biochar additions of 20-25 t ha⁻¹ in temperate forests (Mitchell et al. 2015, 2016; Page-Dumroese et al. 2017). Many incubation experiments have also indicated biochar to affect microbial biomass only at high addition rates (Kolb et al. 2009).

432

Biochar has often been found to increase soil pH especially in acidic soils (Biederman and Harpole 2013). We found that the addition of 10 t ha⁻¹ biochar increased soil pH but the biochar amount of 5 t ha⁻¹ had no effect. Similarly, Noyce et al. (2015) found that the addition of 5 t ha⁻¹ biochar did not affect significantly pH in temperate forest soils. Although biochar was applied on the soil surface, it already had detectable effect on pH in top mineral soil in the second year after treatment in the higher application rate.

439

Although mean N mineralization rates in the organic layer were greater in biochar-amended soils compared with controls, the data showed large variation and differences between treatments were not statistically significant (P > 0.05). Biochar has been shown to increase soil N immobilization in some studies (Bruun et al. 2012; Dempster et al. 2012; Zheng et al. 2013; Ameloot et al. 2015), whereas in 444 some studies biochar has increased nitrification and ammonification (Anderson et al. 2011; Nelissen et al. 2012; Case et al. 2015). Divergent, positive, neutral or negative effects of biochar on N 445 mineralization in literature may exist due to the C:N-ratio of biochar and the C and N status of the 446 447 soil microbes (Prommer et al. 2014). C-rich and N-poor wood biochars may promote N limitation, 448 leading to the retention of produced ammonium in the N-limited microflora, which therefore results in a decrease in the amount of ammonium released to the soil. Conversely, N-rich biochars with low 449 C:N-ratios such as manure-biochars promote microbial C limitation, causing the excess of N to be 450 451 mineralized and therefore N mineralization rates to increase (Prommer et al. 2014).

452

453 Several studies have shown that charcoal from wildfires increases nitrification (Berglund et al. 2004; 454 DeLuca et al. 2002; DeLuca et al. 2006, Ball et al. 2010) likely due to increased soil pH and sorption of phenolic compounds that inhibit nitrifiers (DeLuca et al. 2006). Contrary to hypothesis, biochar 455 456 amendment did not change nitrification rates statistically significantly although soil pH increased. 457 Possibly the increase in pH was too small to affect the nitrification positively. Liming experiments in 458 Finland have shown that the pH increase from 4.1 to 4.4 had little effect on N mineralization 459 (Smolander et al. 1995). Biochar has been even found to decrease nitrification in some studies and it is suggested that volatile organic compounds (VOC's) contained in biochar or increased ethylene 460 emissions after biochar addition, inhibit nitrifiers (Clough et al. 2010; Spokas et al. 2010). Biochar 461 462 may also limit the nitrifier community by reducing the substrate availability by N adsorption to 463 biochar surfaces (Laird et al. 2010) and by microbial N immobilization (Kolb et al. 2009). The observed unchanged nitrification rates suggest that biochar addition does not increase the risk of soil 464 465 N losses through nitrate leaching or gaseous losses through denitrification in the studied ecosystem.

466

467 Biological N fixation has been reported to be 0.1-4 kg N ha⁻¹ yr⁻¹ in boreal forests (Cleveland et al. 468 1999; DeLuca et al. 2002, 2008; Zackrisson et al. 2004; Palviainen et al. 2017b). Our results were at

the lower end of this range $(0.43-0.58 \text{ kg N ha}^{-1} \text{ yr}^{-1})$ which may be a consequence of the early 469 470 successional stage of the investigated forest stands. The rotation period is 90-100 years, and fire return 471 interval is 50-200 years in these types of forests (Ohlson et al. 2009). The biological N fixation rates is estimated to be < 0.5 kg ha⁻¹ yr⁻¹ in early successional forests (Zackrisson et al. 2004; DeLuca et 472 al. 2007). Furthermore, the rather high N deposition (7.4 kg ha⁻¹ yr⁻¹) in our study area (Korhonen et 473 al. 2013) may be one reason for low N fixation rates. N additions of as small as 3 kg N ha⁻¹ yr⁻¹ have 474 475 already shown to lower N fixation in mosses (Gundale et al. 2011). The addition of biochar did not 476 have a significant effect on the biomass of mosses although in several studies biochar has been found 477 to increase the growth of crops and trees (Robertson et al. 2012; Biederman and Harpole 2013; Thomas and Gale 2015). Our results support the findings of Gundale et al. (2015), who did not find 478 10 t ha⁻¹ biochar addition to affect the coverage of mosses in boreal forests. Mosses do not get 479 advantage for biochar induced improved water holding capacity, cation exchange capacity and 480 nutrient availability to a similar extent as vascular plants, because boreal mosses are rather drought-481 482 tolerant and receive the majority of their nutrients from rainwater (Brown and Bates 1990).

483

484 To our knowledge, this is the first study to examine the impacts of biochar amendment on biological N-fixation in boreal forests. Biochar treatments did not have a significant effect on N-fixation which 485 486 is likely due to that soil microbial biomass and moss biomass did not markedly change after biochar 487 addition. In contrast, biochar has been commonly reported to increase N-fixation in leguminous plants in agro-ecosystems and it has been attributed to elevated soil pH and improved nutrient availability 488 489 (Rondon et al. 2007; Mia et al. 2014; Güereña et al. 2015; Van Zwieten et al. 2015). In our study, 490 biochar increased soil pH which may have had positive effect on N-fixation but on the other hand 491 part of the biochar contained N may have been mineralized and this may have affected negatively the 492 N-fixation. Also, Robertson et al. (2012) found that biochar amendment did not change N-fixation 493 rates in the root nodules of alder seedlings. Nitrogen fixation rates increased with temperature, which 494 is consistent with previous findings that N fixation in feather mosses peaks at temperatures of 13°C–
495 22°C, and declines above 30°C temperatures (Gentili et al. 2005).

496

This study explores short-term responses of biochar amendment in a typical boreal forest. We conclude that not all potential impacts are evident just one year after biochar application and hence, specific questions may require long-term experiments. Although our study covered a short response period relative to a typical forest rotation length, it is an important first step in evaluating the impacts of potential biochar application in boreal forests on the C and N cycles. Studies like this, in combination with additional long-term studies, are necessary before biochar use can be promoted and included in C trading schemes in the boreal region.

504

505 Conclusions

506

The results indicate that wood-derived biochar amendment of 5–10 t ha⁻¹ did not have a clear and 507 508 consistent effect on soil CO₂ effluxes in boreal Scots pine forests. Biochar amendment increased the 509 soil pH but it had no significant effect on soil microbial biomass and biological N fixation at this 510 stage. Nitrogen mineralization rates in the organic layer had a tendency to increase with the amount of added biochar, but no statistically significant effect was detected. The results suggest that biochar 511 512 can be utilized to climate change mitigation and C sequestration in boreal forests without causing 513 undesirable effects on soil microbial biomass, key N cycling processes or native soil C stocks. More 514 long-term field studies from forest ecosystems are, however, needed to confirm these perceptions and 515 to find optimum biochar application strategies.

516

517 Acknowledgements

519	This study was funded by The Foundation for Research of Natural Resources in Finland (2016085).
520	The study was also supported by the Academy of Finland (286685, 294600, 307222, 277623) and the
521	FCoE of atmospheric sciences (Center of Excellence (1118615). We thank for the staff of Hyytiälä
522	Forestry Field Station for supporting us in the field work and Marjut Wallner for help with laboratory
523	analyses.
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741	

Table 1. Mean (±SE) soil carbon and nitrogen concentrations, carbon and nitrogen ratio, soil carbon
 and nitrogen pools and soil particle size distribution in the study site.

Soil layer	C (%)	N (%)	C:N ratio	C g m ²	N g m ²	Clay (%)	Silt (%)	Sand (%)
Organic layer	29.82	0.91	33 (0.5)	672	20			
	(1.45)	(0.04)		(53)	(1)			
Mineral soil 0-	3.23	0.11	29 (0.8)	737	25	0.00	15.48	84.52
5 cm	(0.32)	(0.01)		(72)	(3)			
Mineral soil 5–	1.28	0.05	26 (8.7)	746	27	0.00	12.67	87.33
15 cm	(0.06)	(0.003)		(39)	(2)			

747 Table 2. Characteristics of the biochars. Values are mean \pm SE.

	Biochar 500°C	Biochar 650°C
C (%)	60.61 (1.93)	60.61 (1.88)
N (%)	0.87 (0.03)	0.29 (0.10)
C:N ratio	70 (4.1)	364 (70.4)
Loss on ignition (LOI)	82.0 (6.4)	87.5 (1.4)
pH	8.77 (0.23)	8.87 (0.07)
Electric conductivity µs cm ⁻¹	1361 (226)	1462 (179)
Al mg g ⁻¹	21.83 (1.68)	20.78 (1.07)
Ca mg g ⁻¹	190.68 (10.34)	154.15 (7.17)
Cu mg g ⁻¹	0.22 (0.01)	0.20 (0.01)
Fe mg g ⁻¹	24.01 (2.09)	21.32 (1.10)
K mg g ⁻¹	29.81 (3.27)	42.87 (6.73)
Mg mg g ⁻¹	17.08 (1.09)	19.32 (1.55)
Mn mg g ⁻¹	3.86 (0.16)	2.95 (0.10)
Na mg g ⁻¹	3.30 (0.96)	2.71 (0.24)
Ni mg g ⁻¹	0.07 (0.01)	0.08 (0.01)
P mg g ⁻¹	10.78 (0.40)	7.41 (0.21)
S mg g ⁻¹	5.61 (0.25)	4.33 (0.12)
Si mg g ⁻¹	2.29 (0.11)	2.19 (0.05)
Zn mg g ⁻¹	0.42 (0.01)	0.34 (0.01)

Table 3. Monthly mean (\pm SE) soil temperatures (°C) at 5 cm depth in different treatments. The same

752	letters indicate no statistically significant differences among treatments.	
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Treatment	May	June	July	August	September
Control	7.0 (0.4) ^a	11.7 (0.4) ^a	14.7 (0.2) ^a	13.8 (0.2) ^a	10.2 (0.3) ^a
Biochar 500°C 5 t ha ⁻¹	$8.4 (0.4)^{a}$	12.3 (0.4) ^a	15.2 (0.2) ^a	14.3 (0.2) ^a	10.5 (0.3) ^a
Biochar 500°C 10 t ha ⁻¹	8.1 (0.4) ^a	12.8 (0.4) ^a	15.6 (0.2) ^a	14.4 (0.2) ^a	10.4 (0.3) ^a
Biochar 650°C 5 t ha ⁻¹	7.3 (0.5) ^a	12.5 (0.4) ^a	15.3 (0.2) ^a	14.4 (0.2) ^a	10.6 (0.3) ^a
Biochar 650°C 10 t ha ⁻¹	7.5 (0.4) ^a	11.5 (0.4) ^a	14.5 (0.2) ^a	13.9 (0.2) ^a	10.5 (0.3) ^a

Table 4. The average biomass (\pm SE) of mosses (kg ha⁻¹) in different treatments.

	Control	Biochar 5 t ha ⁻¹	Biochar 10 t ha ⁻¹
Pleurozium schreberi	2588 (497) ^{ab}	4313 (588) ^a	2451 (593) ^b
Dicranum polysetum	1978 (801) ^a	1000 (557) ^a	1524 (587) ^a
Total biomass of mosses	4566 (455) ^a	5313 (572) ^a	3975 (582) ^a

760 Figure captions

Figure 1. a) The grain size of the biochar was 5-10 mm b) Biochar was spread manually on the soil surface c) Study site d) The mosses grew on top of the biochar layer one year after biochar addition.

Figure 2. Mean (\pm SE) soil CO₂ effluxes in different treatments in June and July. Different letters indicate statistically significant differences (P <0.05) between treatments.

Figure 3. Mean (\pm SE) soil microbial biomass carbon, soil microbial biomass nitrogen and soil microbial biomass carbon:nitrogen ratio in control plots and 5 t ha⁻¹ and 10 t ha⁻¹ biochar (pyrolysis temperature 650°C) treatments. Different letters indicate statistically significant differences (P <0.05) between treatments.

- **Figure 4.** The mean (\pm SE) biological nitrogen fixation rates (μ g m⁻² d⁻¹) in each incubation temperature (10°C, 15°C and 20°C) in control plots and 5 t ha⁻¹ and 10 t ha⁻¹ biochar (pyrolysis temperature 650°C) treatments. Statistically significant differences (P <0.05) between treatments in each temperature group are indicated by different lower-case letters, whereas statistically significant differences (P <0.05) between temperatures are indicated by upper-case letters in parenthesis.
- **Figure 5.** The mean (\pm SE) soil pH in the organic layer and upper 10 cm mineral soil layer in control plots and 5 t ha⁻¹ and 10 t ha⁻¹ biochar (pyrolysis temperature 650°C) treatments. Different letters indicate statistically significant differences (P <0.05) between treatments.

Figure 6. The mean (\pm SE) net nitrogen mineralization, ammonification and nitrification rates in the organic layer and upper 10 cm mineral soil layer in control plots and 5 t ha⁻¹ and 10 t ha⁻¹ biochar (pyrolysis temperature 650°C) treatments. Different letters indicate statistically significant differences (P <0.05) between treatments.

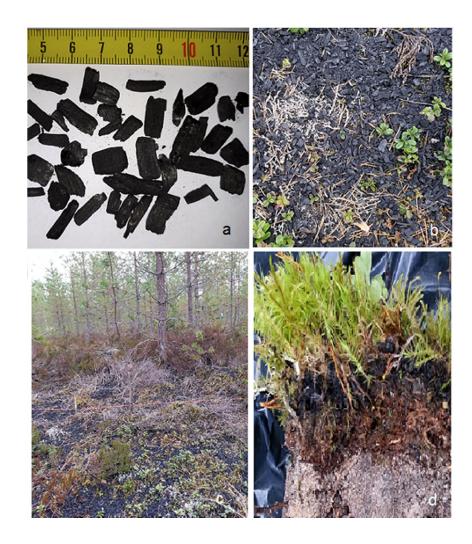
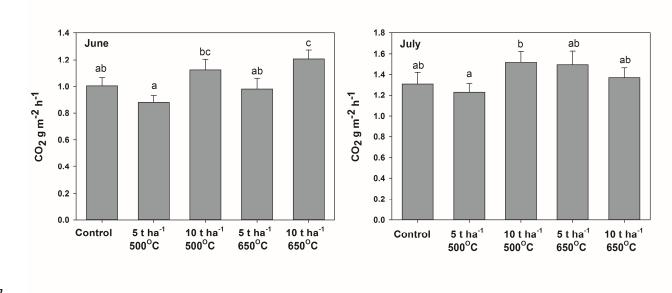


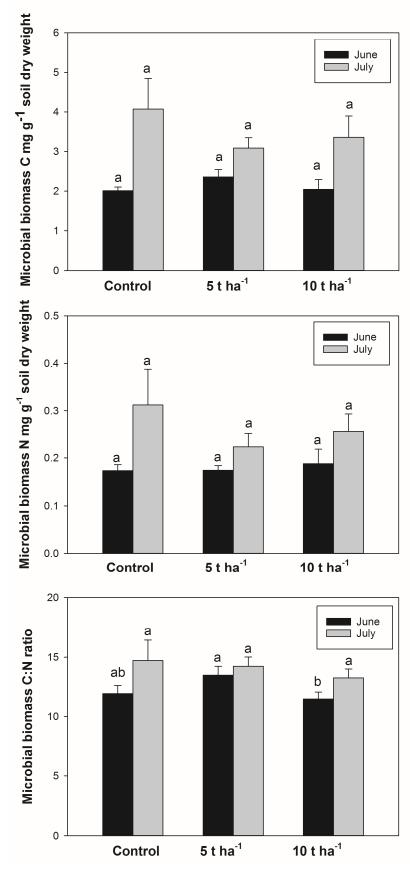


Fig. 1

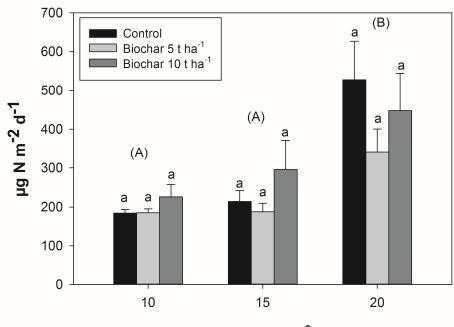




788 Fig. 2



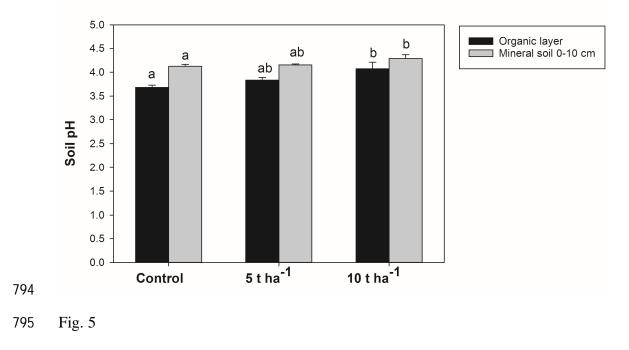
791 Fig.3

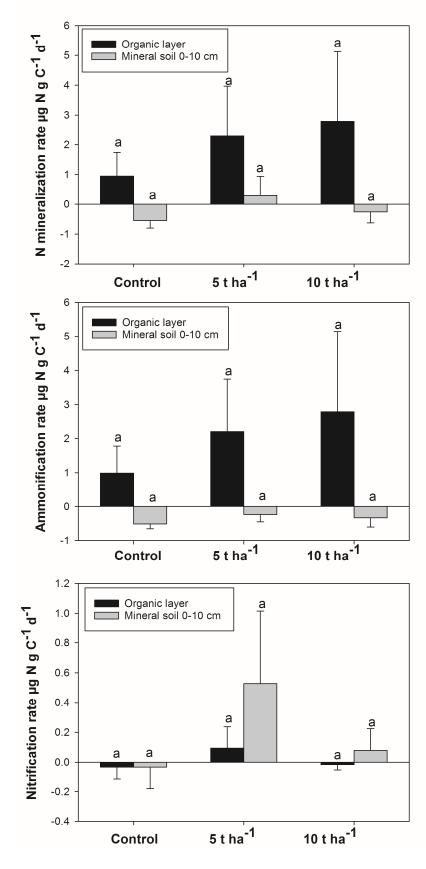


Temperature °C



793 Fig. 4





798 Fig. 6